

IODP Expedition 335: Deep Sampling in ODP Hole 1256D

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Abstract

Observations of the gabbroic layers of untectonized ocean crust are essential to test theoretical models of the accretion of new crust at mid-ocean ridges. Integrated Ocean Drilling Program (IODP) Expedition 335 (“Superfast Spreading Rate Crust 4”) returned to Ocean Drilling Program (ODP) Hole 1256D with the intention of deepening this reference penetration of intact ocean crust a significant distance (~350 m) into cumulate gabbros. Three earlier cruises to Hole 1256D (ODP 206, IODP 309/312) have drilled through the sediments, lavas, and dikes and 100 m into a complex dike-gabbro transition zone.

Operations on IODP Expedition 335 proved challenging throughout, with almost three weeks spent re-opening and securing unstable sections of the hole. When coring commenced, the comprehensive destruction of the coring bit required further remedial operations to remove junk and huge volumes of accumulated drill cuttings. Hole-cleaning operations using junk baskets were successful, and they recovered large irregular samples that document a hitherto unseen sequence of evolving geological conditions and the intimate coupling between temporally and spatially intercalated intrusive, hydrothermal, contact-metamorphic, partial melting, and retrogressive processes.

Hole 1256D is now clean of junk, and it has been thoroughly cleared of the drill cuttings that hampered op-

erations during this and previous expeditions. At the end of Expedition 335, we briefly resumed coring before undertaking cementing operations to secure problematic intervals. To ensure the greatest scientific return from the huge efforts to stabilize this primary ocean lithosphere reference site, it would be prudent to resume the deepening of Hole 1256D in the nearest possible future while it is open to full depth.

Why Study Ocean Crust Forming at Fast Spreading Rates?

The vast majority (~70%) of magma derived from the mantle is brought into the Earth’s crust at the mid-ocean ridges, and approximately two-thirds of that magma cools

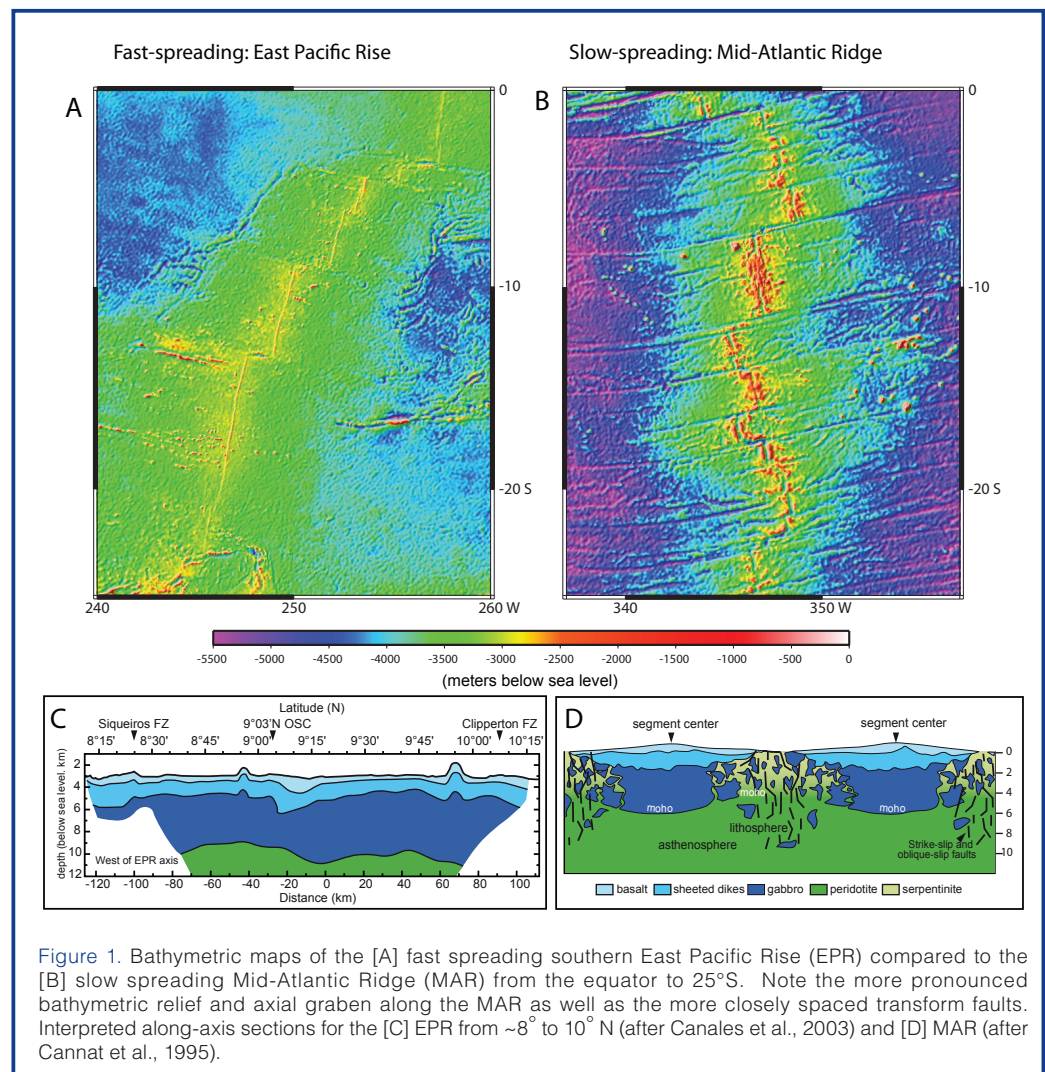


Figure 1. Bathymetric maps of the [A] fast spreading southern East Pacific Rise (EPR) compared to the [B] slow spreading Mid-Atlantic Ridge (MAR) from the equator to 25°S. Note the more pronounced bathymetric relief and axial graben along the MAR as well as the more closely spaced transform faults. Interpreted along-axis sections for the [C] EPR from ~8° to 10° N (after Canales et al., 2003) and [D] MAR (after Cannat et al., 1995).

and crystallizes in the lower portion of the ocean crust. Seismic, bathymetric, and marine geological observations indicate that ocean crust formed at fast spreading rates (full rate $>80 \text{ mm yr}^{-1}$) is much less variable than crust formed at slow spreading rates ($<40 \text{ mm yr}^{-1}$; Fig. 1) and is closer to the ideal “Penrose” pseudo-stratigraphy developed from ophiolites. Hence, understanding fast-spreading accretion processes at a few sites might reasonably be extrapolated to describe a significant portion of the Earth’s surface. Although $<20\%$ of modern ridges are moving apart at fast spreading rates (Fig. 2), nearly 50% of present-day ocean crust and $\sim 30\%$ of the Earth’s surface were produced by this pace of spreading. The great majority of crust that subducted back into the mantle over the past $\sim 200 \text{ m.y.}$ formed at fast-spreading ridges (Müller et al., 2008), making characterizing this style of crust most relevant for understanding the recycling of crustal and ocean-derived components back into the mantle.

The rate of spreading of the oceanic lithosphere has profound effects on the style of accretion at mid-ocean ridges due to changing balances between plate motion, magma production, conductive and hydrothermal cooling, detachment tectonics, and serpentinization of the upper mantle (Fig. 1c vs. 1d; Dick, 1989; Cannat et al., 1995, 2004, 2009; Chen and Phipps Morgan, 1996; Dick et al., 2003; Escartin et al., 2008). Although insights on formation of intrusive crust at detachment-dominated, slow-spread lithosphere have been obtained (e.g., Holes 735B, U1309D, and other sites; ODP Legs 118, 153, 176, 209, and IODP Expeditions 304–305; Dick et al., 2000; Ildefonse et al., 2007; Kelemen et al., 2007), the thermal regime and the melt supply and delivery in these settings differ significantly from those of the axial zone in fast-spreading lithosphere. Detailed understanding of the relatively uniform mechanisms operating at fast spreading ridges would provide a vital benchmark against which heterogeneous accretion on slow-spreading ridges could be compared.

Geological Setting of Site 1256 and Science Objectives of IODP Expedition 335

IODP Expedition 335 (13 April to 3 June 2011) was the fourth scientific drilling cruise of the Superfast campaign, and it returned to ODP Hole 1256D ($6^{\circ}44.163'N, 91^{\circ}56.061'W$) to deepen this ocean crust reference penetration a significant distance into cumulate gabbros. ODP Hole 1256D is located on 15-Ma-crust in the eastern equatorial Pacific Ocean in oceanic basement that formed during a sustained episode of superfast ocean ridge spreading ($>200 \text{ mm yr}^{-1}$; Wilson, 1996). Ocean crust formed at a superfast spreading rate was deliberately targeted because there is strong evidence from mid-ocean ridge seismic experiments that gabbros occur at shallower depths in intact ocean crust with higher spreading rates. Therefore, the often difficult-to-drill upper ocean crust should be relatively thin. Expedition 335 follows on from ODP Leg 206 in 2002 and IODP Expeditions

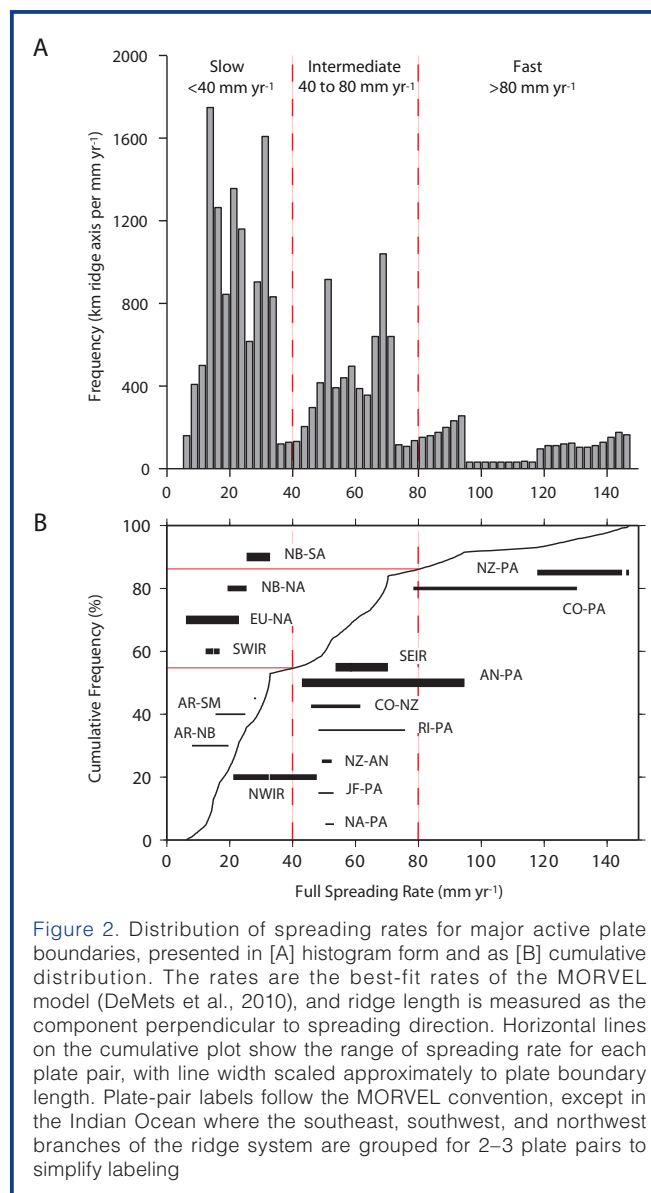


Figure 2. Distribution of spreading rates for major active plate boundaries, presented in [A] histogram form and as [B] cumulative distribution. The rates are the best-fit rates of the MORVEL model (DeMets et al., 2010), and ridge length is measured as the component perpendicular to spreading direction. Horizontal lines on the cumulative plot show the range of spreading rate for each plate pair, with line width scaled approximately to plate boundary length. Plate-pair labels follow the MORVEL convention, except in the Indian Ocean where the southeast, southwest, and northwest branches of the ridge system are grouped for 2–3 plate pairs to simplify labeling

309/312 in 2005 (Wilson et al., 2003; Teagle et al., 2006; Wilson et al., 2006) that prepared the first scientific borehole for deep drilling by installing a large reentry cone secured with almost 270 meters of 16-inch casing through the 250-m-thick sedimentary overburden and cemented into the uppermost basement. Hole 1256D was then deepened through a $\sim 810\text{-m}$ -thick sequence of lavas and a thin ($\sim 346 \text{ m}$) sheeted dike complex, the lower 60 meters of which is strongly contact metamorphosed to granoblastic textures (Fig. 4). The first gabbroic rocks were encountered at 1407 meters below seafloor (mbsf) where the hole entered a complex dike-gabbro transition zone that includes two 20- to 50-m thick gabbro lenses intruded into granoblastic dikes. As of the end of Expedition 312 in December 2005, Hole 1256D had a total depth of 1507.1 mbsf, and it was open to its full depth.

It was expected that even with a shortened cruise (~ 6 weeks), coring on IODP Expedition 335 would be able to deepen Hole 1256D $\sim 350 \text{ m}$, completely through the dike-gabbro transition zone and MOR textured gabbros and a

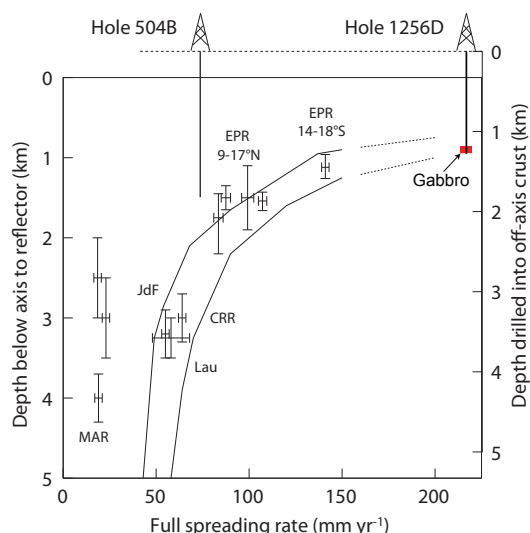


Figure 3. Depth to axial low-velocity zone plotted against spreading rate (Purdy et al., 1992; Carbotte et al., 1997). Depth versus spreading rate predictions from two models of Phipps Morgan and Chen (1993) are shown, extrapolated subjectively to 200 mm yr⁻¹. Penetration to date in Holes 504B and 1256D is shown, with the depth at which gabbros were intersected indicated by the red box. Following core descriptions, a thickness of ~300 m of off-axis lava is shown for Hole 1256D and assumed for Hole 504B. EPR=East Pacific Rise, JdF=Juan de Fuca Ridge, Lau=Valu Fa Ridge in Lau Basin, CRR=Costa Rica Rift, MAR=Mid-Atlantic Ridge.

significant distance into true cumulate gabbros. Listed below are the specific scientific objectives of IODP Expedition 335 that were to be addressed by deepening Hole 1256D a significant distance into cumulate gabbros.

What is the major mechanism of magmatic accretion in crust formed at fast spreading rates? Is the lower crust formed by gabbro glaciers or sheeted sills (Korenaga and Kelemen, 1998; Fig. 5) or by some mixed or unknown mechanism?

How is heat extracted from the lower ocean crust? Specifically, what is the role of hydrothermal circulation in extracting latent heat from the lower ocean crust?

What is the geological significance of the seismic layer 2/3 boundary at Site 1256?

What is the magnetic contribution of the gabbro layer? Can the magnetic polarity structure of the lower crust be used to constrain cooling rates?

Operations

Expedition 335 reentered Hole 1256D more than five years after the last expedition to this site. The expedition encountered and overcame a series of significant engineering challenges, including a major obstruction at a depth of ~922 mbsf that had initially prevented reentry into the hole to its full depth of 1507 meters. At the bottom of the hole, coring resumed in the very hard granular rocks recovered during the previous Superfast expedition. Although there may only be a few tens of meters of these particularly tenacious granoblastic basalts formed by the contact metamorphism of sheeted dikes, their extreme toughness once more proved challenging to sample, resulting in the grinding down of one of the hardest formation coring bits (C9) into a smooth stump (Fig. 6a,b).

Action was then undertaken to clear the bottom of the hole of metal debris from the failed coring bit and drilling cuttings. These efforts returned hundreds of kilograms of rocks and drill cuttings (Fig. 6d), including large blocks (up to 3.5 kg) of the

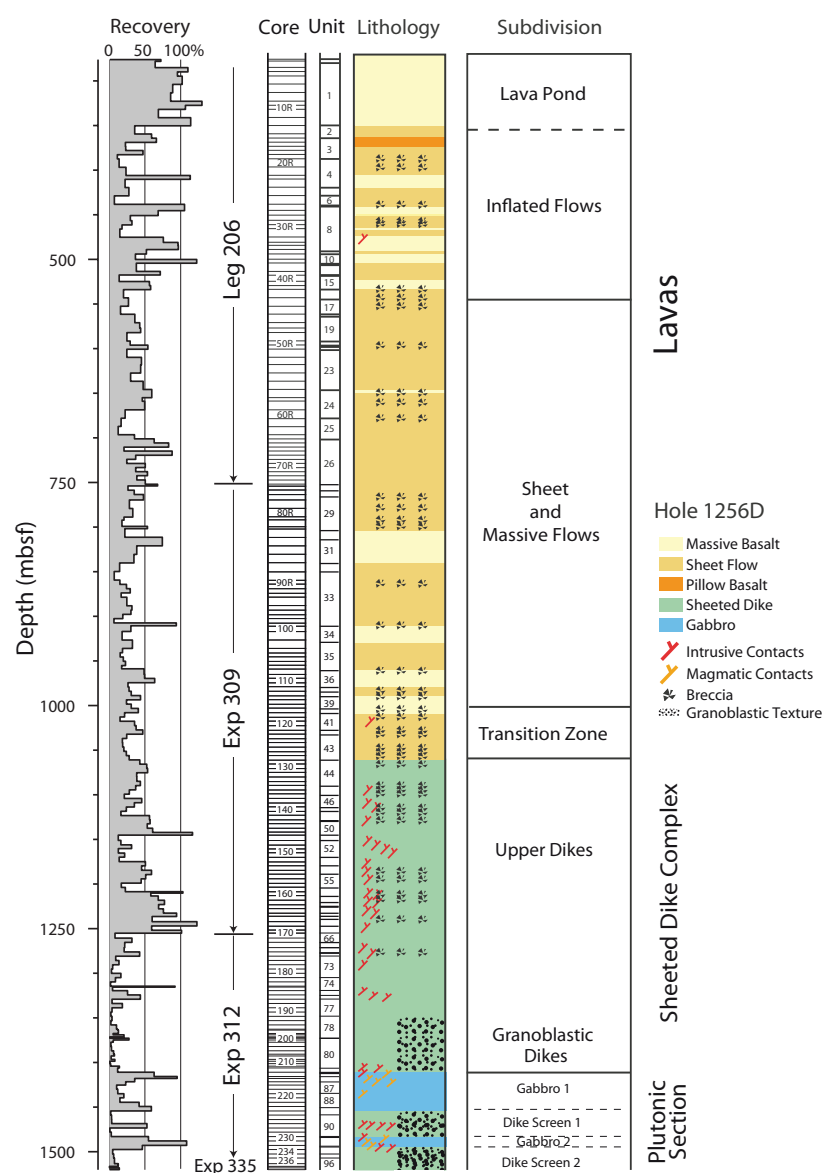


Figure 4. Stratigraphic column for Hole 1256D at the end of Expedition 335, showing the major and minor lithologic divisions of the upper ocean crust.

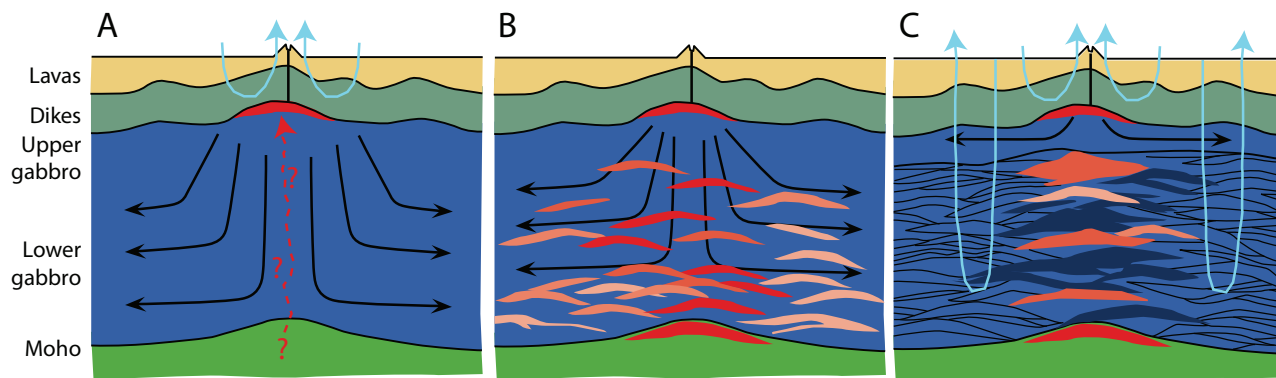


Figure 5. Diagrams of crustal accretion models (modified from Korenaga and Kelemen, 1998). Black arrows show the movement of the solid lower crust; blue arrows show the dominant zones where hydrothermal circulation will remove latent and sensible heat; red arrows show the movement of magma (unknown in all models). [A] Gabbro glacier ductile flow model (Henstock et al., 1993; Phipps Morgan and Chen, 1993; Quick and Denlinger, 1993). Ductile flow down and outward from a high level axial magma chamber constructs the lower crust. [B] A hybrid model of ductile flow with sill intrusions (Boudier et al., 1996). [C] "Sheeted" or "stacked" sill model of *in situ* formation of the lower crust by on-axis sill intrusions (Bédard et al., 1988; Kelemen et al., 1997; Kelemen and Aharonov, 1998; MacLeod and Youancq, 2000). (Original figure by Rosalind Coggon).

culprit granoblastic basalts that hitherto had only been very poorly recovered through coring. A limited number of gabbro boulders were also recovered, indicating again that we are tantalizingly close to breaking through into the gabbroic layer.

Scientific Results

Hole opening and remediation operations resulted in a major loss of time from the coring and wireline activities planned for Expedition 335. Coring on this expedition deepened Hole 1256D only modestly, from 1507.1 mbsf to 1521.6 mbsf (Fig. 4) at low rates of penetration (0.9 m hr^{-1}) and recovery (11%).

Hole cleaning operations at the bottom of Hole 1256D, particularly those run that deployed the reverse-circulation junk basket, brought back a unique collection of large cobbles, angular rubble, and fine cuttings of principally strongly to completely recrystallized granoblastic basalt with minor gabbroic rocks and evolved plutonic rocks. The granoblastic basalts record high-temperature (granulite facies) metamorphism of previously hydrothermally altered sheeted dikes, with unequivocal evidence for hydrothermal alteration before and after the formation of granoblastic textures (Teagle et al., 2006; Koepke et al., 2008; France et al., 2009; Alt et al., 2010). The large blocks exhibit intrusive, structural, and textural relationships, and overprinting and cross-cutting hydrothermal alteration and metamorphic paragenetic sequences. Some of the rubble and ~30% of the fine cuttings recovered by fishing and cleaning operations most likely came from the lava sequences at the top of the hole on the basis of igneous textures and low-temperature alteration minerals (Mg-saponite, amorphous silica). In contrast, the high extent of metamorphic recrystallization exhibited by the granoblastic basalts, and operational factors (e.g., pipe movements), provide strong evidence that the granoblastic

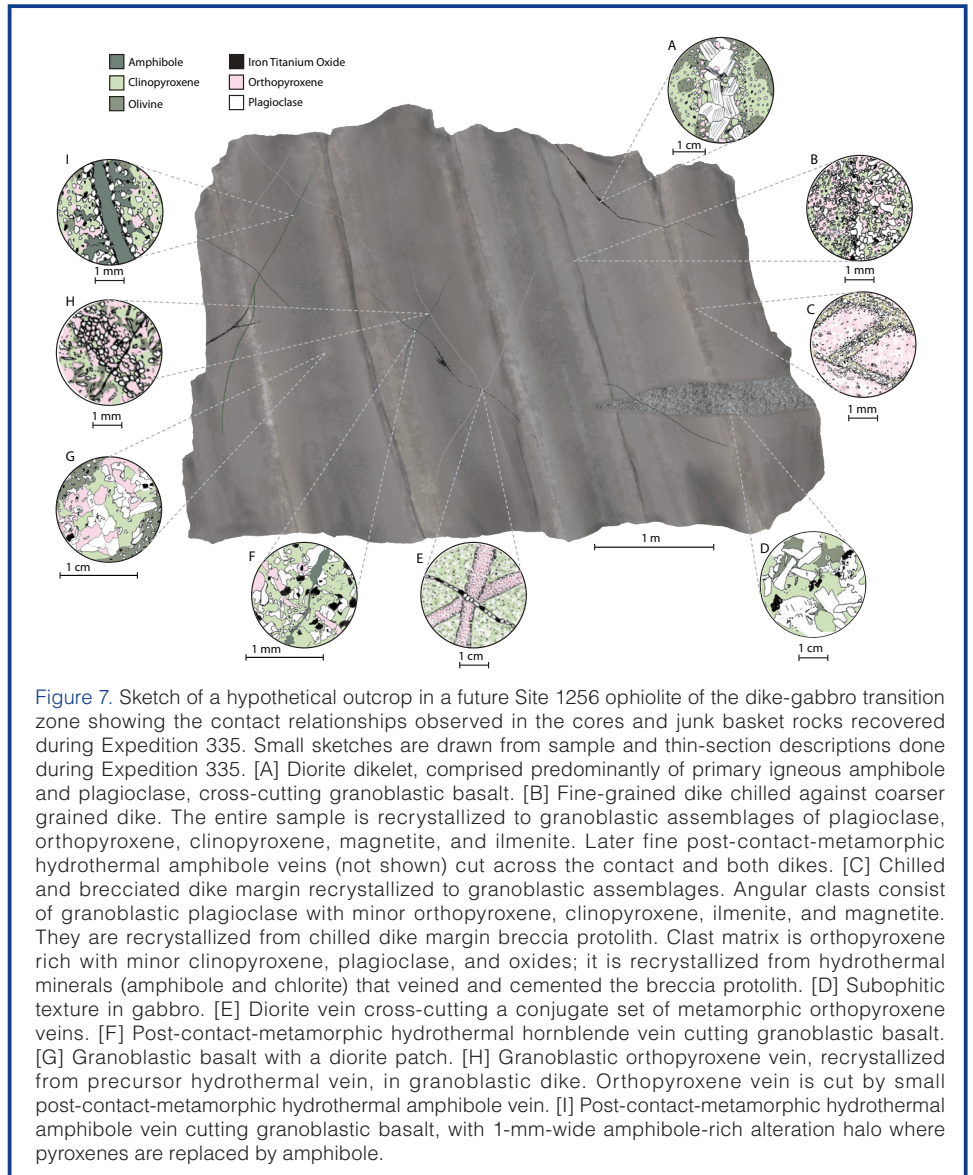


Figure 6. Photographs of some drilling, fishing, and milling tools illustrating the multiple operations achieved during IODP Expedition 335. [A] & [B] Remains of the Ultrera C9 RCB coring bit used during Run 9. The bit was probably used for ~10 h after it disintegrated, which resulted in this spectacularly abraded and sculptured bit ("Stumpy"), something never seen before by the drillers. [C] Bottom of the Bowen Reverse Circulation Junk Basket (RCJB), showing its hard-facing structure and the junk catcher spring fingers inside. This tool recovered a series of large-scale cobbles (up to 5 kg) of granoblastic basalt. [D] A series of RCJB runs returned a bottom hole assembly packed by fine-grained cuttings. This contributed efficiently to cleaning Hole 1256D. [E] Heavily worn and under-gauge 9-inch flat-bottomed milling tool. This tool worked at the bottom of the hole for 12 h; its final state indicates the very abrasive nature of metal debris and/or rocks at the bottom of the hole and an under-gauge lowermost portion (tens of centimeters) of the hole. Note for comparison the hard-facing structure of the next milling tool on the right side of the picture. [F] 9-inch Smith FH3VPS tricone bit used for Run 16. This armored bit was efficient to ream and clean the under-gauge bottom of the hole.

basalts, minor gabbros, and evolved plutonic rocks were sourced from the lowermost reaches of Hole 1256D (1494.9–1521.6 mbsf), most probably from below the lower-most recovered gabbroic interval (Fig. 4). Hence, the rocks recovered during Expedition 335 represent a ~15-m interval of the upper crust/lower crust transition, occurring below the ~90-m section, recovered during Expedition 312, of two gabbroic bodies (Gabbro 1 and Gabbro 2) separated by Dike Screen 1 (Fig. 4).

When the textural and contact relationships exhibited by the rocks recovered from the junk baskets are placed in the geological context of the Hole 1256D stratigraphy, a vision emerges of a complex, dynamic thermal boundary layer zone. This region of the crust between the principally hydrothermal domain of the upper crust and the intrusive magmatic domain of the lower crust is one of evolving geological conditions. There is intimate coupling between temporally and spatially intercalated magmatic, hydrothermal, partial melting, intrusive, metamorphic, and retro-grade processes.

With only a minor depth advance in Hole 1256D, we are yet to recover samples of cumulate gabbros required to test models of ocean ridge magmatic accretion and the intensity of hydrothermal cooling at depth. The total vertical thickness of granoblastic basalts is over 114 m, and Dike Screen 2 is now about the same thickness (so far) as Dike Screen 1. Highly perched, isolated gabbro intrusions are uncommon in ophiolites. The energy requirements for the granoblastic recrystallization at granulite facies condition of a >114-m-thick zone of sheeted dikes massively exceeds the thermal capacity of Gabbros 1 and 2 (Coggon et al., 2008; Koepke et al., 2008) if a simple sub-horizontal arrangement of these layers is assumed. The enormous heat requirements for such extensive granulite facies recrystallization, the evidence for partial melting, together with the tantalizing presence of minor but not uncommon gabbroic rocks and felsic intrusives, dikelets, and veins, strongly indicates that the layer of purely plutonic rocks should be at most only a few tens of meters deeper in the hole.



Conclusions

Although the extensive remedial operations on Expedition 335 precluded significant deepening of Hole 1256D, significant progress was made in improving the stability of the borehole. The most problematic out-of-gauge zone at ~920–960 mbsf that caused reentry problems on Expeditions 312 and 335 has been stabilized with cement. The bottom of the hole has been cleared of rubble and junk, and there appears to be only a short (<1 m) under-gauge zone. What is important is that the regular, large sweeps of high viscosity mud have finally overcome and expunged the vast amount of fine cuttings recirculating in the hole, some most likely resident since ODP Leg 206. This is shown by the absence of soft fill in the final ~5 reentries, compared to upwards of 50 meters of soft fill at the end of Expedition 312. The engineering efforts on Expedition 335 have repaired and prepared Hole 1256D for further deep drilling, following five years of neglect. Hole 1256D is 1500 m of hard rock coring closer to cumulate gabbros than any other options in intact ocean crust and once more poised to answer fundamental

questions about the formation of new crust at fast spreading mid-ocean ridges. This would be best achieved by a timely return to the site.

Acknowledgments

IODP Expedition 335 was not an easy cruise. Engineering difficulties throughout the expedition presented a continuous series of challenges. The Shipboard Science Party is extremely grateful for the immense efforts of the transocean crew aboard *JOIDES Resolution* for their perseverance, innovation, open discourse, and “can-do” attitude. This team was aboard when Hole 1256D was first spudded in and cased into basement on ODP Leg 206; they were not going to allow the hole to fail. The Texas A&M University (TAMU) Operations Superintendent Ron Grout has sailed on three of the four cruises to Hole 1256D, and his advice in testing situations proved invaluable. The proponents of the Superfast campaign thank the IODP Science Advisory Structure (SAS) for their extension of Expedition 335; without the extra two weeks there could have been the catastrophic loss of Hole 1256D, but this extra time allowed the clearing of junk and cuttings and preservation of the hole for future deepening into the cumulate gabbro layers.

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Photo Credit

Fig. 6: Photos [A] to [D], and [F]: Benoit Ildefonse, CNRS; Photo [E]: Johan Lissenberg, Cardiff University.