

COBBOOM: The Continental Breakup and Birth of Oceans Mission

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Introduction

The rupture of continents and creation of new oceans is a fundamental yet primitively understood aspect of the plate tectonic cycle. Building upon past achievements by ocean drilling and geophysical and geologic studies, we propose “The Continental Breakup and Birth of Oceans Mission (COBBOOM)” as the next major phase of discovery, for which sampling by drilling will be essential.

In September 2006, fifty-one scientists from six continents gathered in Pontresina, Switzerland to discuss current knowledge of continental breakup and sedimentary basin formation and how the Integrated Ocean Drilling Program (IODP) can deepen that knowledge (Coffin et al., 2006). Workshop participants discussed a global array of rifted margins (Fig. 1), formulated the critical problems to be addressed by future drilling and related investigations, and identified key rift systems poised for IODP investigations.

Past Achievements

Scientific ocean drilling has played an essential role in the exploration of rifted continental margins. The North Atlantic Rifted Margins (NARM) endeavor of the Ocean Drilling Program (ODP) addressed conjugate margin pair rift systems ranging from “magma-dominated” (Norway/British Isles-Greenland margins of the offshore-onshore early Tertiary North Atlantic large igneous province (LIP)) to “magma-starved” (Iberia-Newfoundland margins of Late Triassic to Early Jurassic age, Fig. 2.) Geophysical studies

and drilling results from these two conjugate pair rift systems have profoundly changed our view of the processes responsible for such margins.

The drilling of ‘seaward dipping reflector’ (SDR) wedges of the North Atlantic LIP off the British Isles (Roberts et al., 1984), Norway (Eldholm et al., 1987, 1989) and SE Greenland (Duncan et al., 1996; Larsen et al., 1994, 1999) confirmed them to be a thick series of subaerial lava flows covering large areas. Lavas on the landward side of the SDRs show geochemical evidence of contamination by continental crust, implying that they rose through continental crust during early rifting, whereas oceanward SDR lavas appear to have formed at a seafloor spreading center resembling Iceland. Drilling results from these margins document extreme magmatic productivity over a distance of at least 2000 km during continental rifting and breakup with temporal and spatial influence of the Iceland plume during rifting, breakup, and early seafloor spreading (Saunders et al., 1998).

Other margins such as Iberia-Newfoundland (Tucholke et al., 2007) appear magma-starved and have been hyperextended by progressive rifting (Lavie and Manatschal, 2006) in stages, with distinct tectonic characteristics controlled by the rheological effects of the gradual thinning of continental crust and uplift of the underlying mantle (Fig. 3). When the crust has been thinned by normal faulting to less than ~10 km, it becomes entirely brittle (Pérez-Gussinyé and Reston, 2001), and tectonism then transitions to spatially focused, closely spaced normal faults that sole into a serpentine detachment at the crust mantle boundary that eventually unroofs upper mantle rocks along a detachment fault exposed at the seafloor. Exhumation of upper mantle rocks continues over a potentially wide region, until mantle uplift generates sufficient magma to initiate seafloor spreading (Tucholke et al., 2007). These results suggest that depth-dependent stretching (DDS) and detachment faulting are major controls on continental rupture and ocean formation.

The ODP also addressed active rifting along a low-angle normal fault in the Woodlark Basin that is propagating into continental lithosphere (Huchon et al., 2002; Taylor et al., 1999a, b).

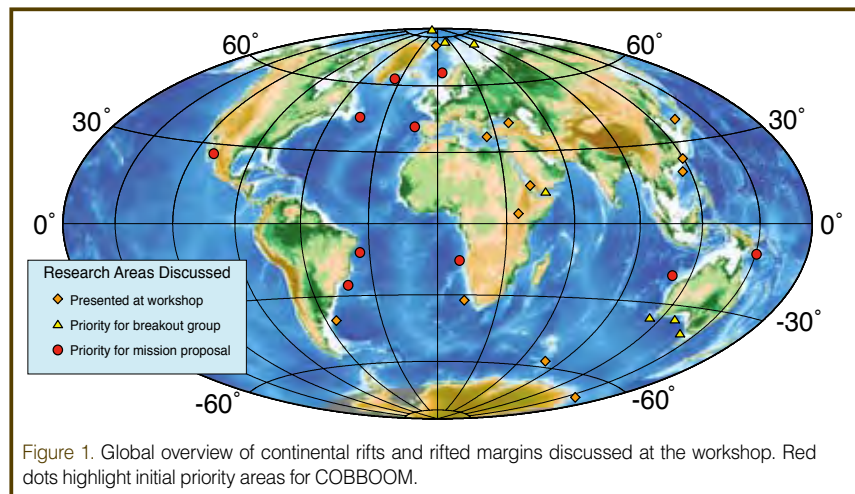
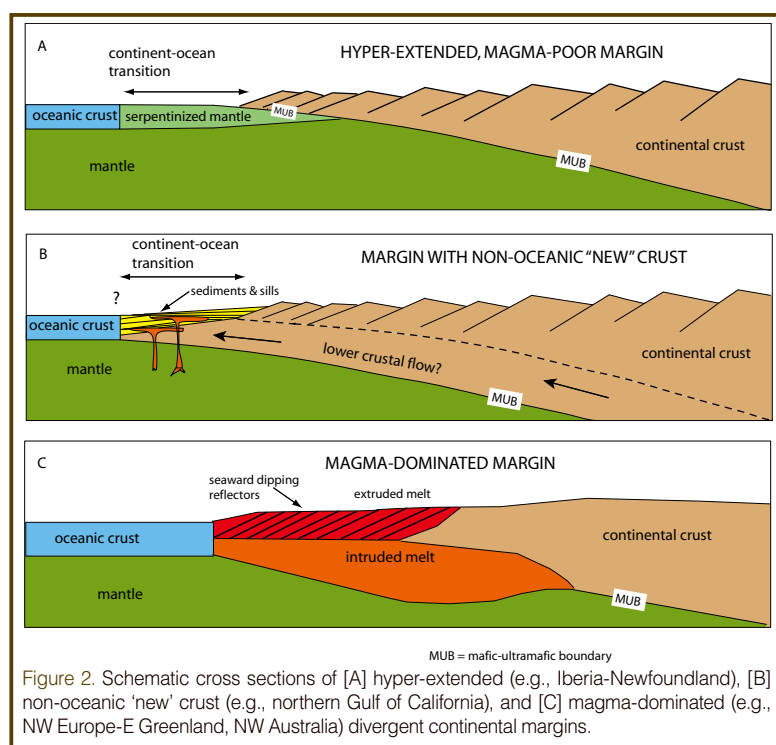


Figure 1. Global overview of continental rifts and rifted margins discussed at the workshop. Red dots highlight initial priority areas for COBBOOM.



Scientific Objectives Associated with Continental Rifting

Variations in the importance and, in particular, the volume of magmatism have led to the classification of margins as “volcanic” or “non-volcanic” (Mutter et al., 1988); however, this binary dichotomy fails to adequately reflect that rifted margins form a spectrum from magma-rich to magma-poor. The key distinction is whether magmatism is more or less than expected from the degree of lithospheric thinning and passive asthenospheric upwelling of normal temperature mantle. Equally important are the timing of magmatism and the strain distribution across margins, i.e., hyper-extended versus a more abrupt transition between continental and oceanic lithosphere. Because a continuum between possible end-members may exist, the focus should be on understanding the fundamental processes causing such variations. Key aspects of continental breakup can only be addressed by drilling and associated studies (Table 1).

More specifically, we need to determine the following at multiple, carefully selected rifted margins: 1) uplift and sub-

sidence history; 2) ages and facies of synrift and syn-faulting sediment; 3) timing, volume, chemistry, and style of magmatism; 4) orientation of deformation fabrics, including faults; and 5) ages and facies of postrift sediment. Such information can be used to infer distribution of strain in space and time; deformation mechanics and dynamics; processes within the mantle, including depth and degree of melting, melt migration, and infiltration; and mantle composition, heterogeneity, and dynamics.

We propose drilling programs on well characterized and representative examples, conjugate where possible, of both active and mature rifted margins ranging from magmatic to amagmatic and abrupt to hyperextended. The rift systems described below constitute an initial focus of investigations for COBBOOM.

Gulf of California

Key Aspects and Problems to be Addressed: The active Gulf of California rift system (Fig. 4) varies along strike in crustal thickness, synrift sediment facies, amount of magmatism, structural style, and width of new seafloor (Lizarralde et al., 2007). The northern basins host an enigmatic type of crust that is 15–20 km thick, characterized by gravity anomalies and seismic velocities suggestive of silicic as opposed to basaltic material (González-Fernández et al., 2005). Low-angle normal faults are also accessible to both onshore sampling and drilling. In the central basins, magma-sediment interactions and fluid/geochemical fluxes, including methanogenesis, will be studied (Fig. 5). In southern segments, the processes and timing of the synrift to postrift transition (breakup unconformity, basin evolution, margin uplift or subsidence) will be examined.

Regional Setting and Background: The system formed from a major reorganization of the Farallon-North American plate boundary during Neogene time (Lonsdale, 1989). Narrow perched basins adjacent to seafloor spreading centers characterize the southernmost segment, whereas diffuse deformation in an apparent continental setting dominates the northern Gulf (Persaud et al., 2003). In the central Gulf, two segments of the Guaymas basin are narrow and

Table 1. Key aspects of continental breakup.

Rift initiation	Driving forces, rift localization, lithospheric strength, thermal structure.
Tectonics of rifting	Distribution of strain, rheological evolution, mechanisms of crustal thinning, strength of faults, 3-D rift geometry, mantle exhumation, transition to seafloor spreading.
Magmatism during rifting	Melt-rift interactions, mantle heterogeneities, melt production into seafloor spreading stage, controls on melt production.
Initiation of seafloor spreading	When and where, development of seafloor spreading magnetic anomalies, mantle thermal structure, mantle sources.
Sedimentary processes and basin evolution	Stratigraphic responses to rifting and breakup, stratigraphy-strain rate relationships, fault patterns and evolution, interactions among erosion, sedimentation, and tectonism.
Environmental consequences & impact	Magma interactions with sediment, hydrosphere, and the atmosphere; tectonic and magmatic controls on ocean gateways.

have slightly thicker crust, suggesting more magmatic input, whereas four segments of the south-central domain are wide, magma-poor rifts (Lizarralde et al., 2007). Simple plate kinematics cannot explain these changes in style because of constant total strain along the rift axis. Rather, along strike variation in pre-existing lithospheric and mantle structure, thermal state, and sediment inputs must have controlled this development.

Proposed Drill Sites: Several thousand kilometers of multichannel seismic lines (Persaud et al., 2003; Lizarralde et al., 2007; González-Fernández et al., 2005; Aragón-Arreola and Martín-Barajas, 2007) image key structures of the rift basins well suited for addressing the following topics: 1) the possible role of lower crustal flow to fill the gap created by rifting; 2) the role of detachment faults in early rifting and/or in delaminating continental crust; 3) differences in magmatism in adjacent rift segments along strike; and 4) the relationships between magmatism and global environmental changes (Svensen et al., 2004; Dickens, 2004).

In the northern Gulf (Fig. 4) high heat flow and thick sediment may have caused lower crustal flow and a diffuse rift with enigmatic transitional crust (González-Fernández et al., 2005). It is feasible that sites can sample possible low-angle normal faults and a complete sedimentary section constraining the age of rifting. Igneous intru-

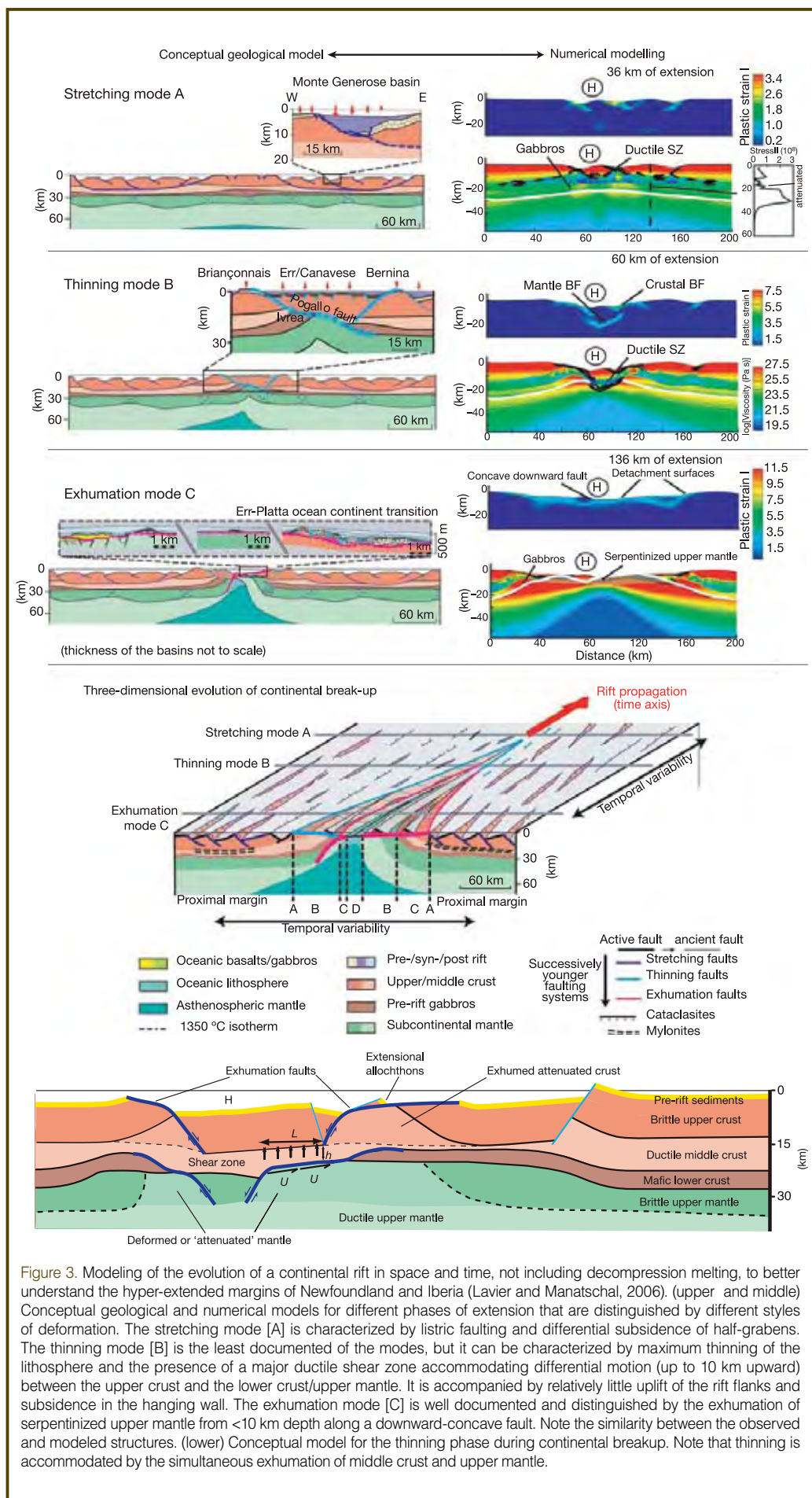


Figure 3. Modeling of the evolution of a continental rift in space and time, not including decompression melting, to better understand the hyper-extended margins of Newfoundland and Iberia (Lavie and Manatschal, 2006). (upper and middle) Conceptual geological and numerical models for different phases of extension that are distinguished by different styles of deformation. The stretching mode [A] is characterized by listric faulting and differential subsidence of half-grabens. The thinning mode [B] is the least documented of the modes, but it can be characterized by maximum thinning of the lithosphere and the presence of a major ductile shear zone accommodating differential motion (up to 10 km upward) between the upper crust and the lower crust/upper mantle. It is accompanied by relatively little uplift of the rift flanks and subsidence in the hanging wall. The exhumation mode [C] is well documented and distinguished by the exhumation of serpentinized upper mantle from <10 km depth along a downward-concave fault. Note the similarity between the observed and modeled structures. (lower) Conceptual model for the thinning phase during continental breakup. Note that thinning is accommodated by the simultaneous exhumation of middle crust and upper mantle.

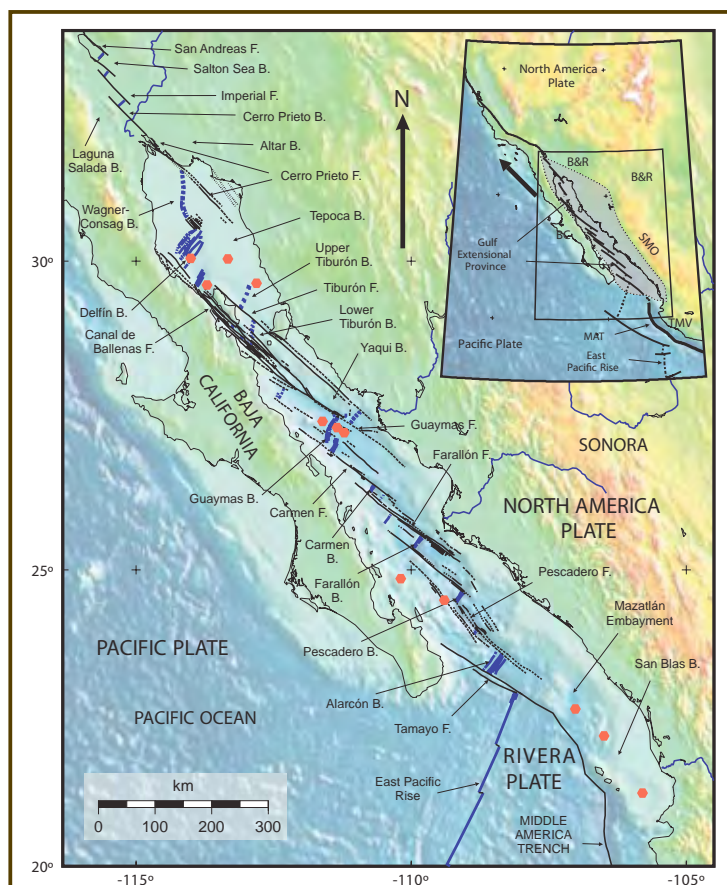


Figure 4. Tectonic map with digital elevations of the Gulf of California region, modified from Aragón-Arreola (2006). This region hosts the modern transensional boundary between the Pacific and North America plates, comprising spreading centers or extensional basins (blue lines) and faults (black lines). Modern and abandoned basins of the plate boundary are labeled. Red dots show proposed IODP sites. Abbreviations: F=fault, B=basin. Inset: regional map of plate boundary system. Inset abbreviations: SMO=Sierra Madre Occidental; B&R = Basin and Range Province; BC=Baja California; TMV=Trans-Mexican Volcanic Belt; MAT=Middle America Trench.

sions may constrain interactions among mantle and lower crustal melts and sediment. Ignimbrites erupted during rifting will provide a detailed volcanic history and chronology for the sedimentary section. Sites in the central Guaymas basin (Fig. 4) will address methanogenesis related to igneous intrusion into organic-rich sediment, a potentially important process in the global carbon cycle that is likely to operate in most sedimented rift systems. Sites in the southern Gulf target magma-poor, tectonic extension during rifting, and the transition from rifting to seafloor spreading. Sites along the Alarcón segment, combined with observations from its landward extension (Umhoefer et al., 2007), will address the extension and subsidence history of this segment as it transitioned from magma-poor, tectonic extension to seafloor spreading.

Complementary to drilling, land studies can constrain the age and composition of rift-related volcanism, and provide additional constraints on basin history.

Woodlark Basin

Key Aspects and Problems to be Addressed: In the Woodlark Basin, westward mid-ocean ridge propagation into continen-

tal crust (Fig. 6) allows detailed investigations of continental lithosphere before, during, and after seafloor spreading commences.

Regional Setting and Background: The Woodlark rift is continuous along strike with a seafloor spreading system (Fig. 6). Since Late Miocene time, continental lithosphere of the Papuan Peninsula thickened during Australia-Pacific plate convergence (Davies and Smith, 1971) and has subsequently rifted at some of the highest known rates (Abers, 2001; Wallace et al., 2004). Seafloor spreading initiated after ~200 km of continental extension in the eastern part of the rift basin prior to 6 Ma, and propagated ~800 km to the modern rift-drift transition adjacent to Moresby Seamount (Taylor et al., 1999a). Adjacent to the westernmost spreading segment is an active, north-dipping low-angle (~30°) normal fault that is currently being dissected by igneous intrusions (Goodliffe and Taylor, 2007). Farther west, seismically active rifts (Ferris et al., 2006) continue toward the active metamorphic core complexes of the D'Entrecasteaux Islands, where exposures of high and ultra-high pressure metamorphic rocks suggest exhumation rates of ~10 km Myr⁻¹ (Baldwin et al., 2004; Monteleone et al., 2007). The upper plates, separated from the lower plates by shear zones and detachment faults (Hill et al., 1992; Little et al., 2007), consist largely of undeformed mafic and ultramafic rocks. Mylonitic lineations and corrugation surfaces parallel Plio-Pleistocene plate motion vectors (Little et al., 2007).

Proposed Drill Sites: Drilling will address two fundamental issues related to the rift-drift transition in easternmost Papua New Guinea.

- 1) Rift-drift transition processes. A drilling transect across a nascent spreading segment will address a) the origin of the first magmas at a new spreading center; b) the relationship between magma supply rate and the development of magnetic anomalies associated with seafloor spreading; c) magma-sediment interactions; d) the state of stress at the rift-drift transition; and e) how plate motion accommodation transitions from a low-angle fault to crustal accretion. Drilling will sample the center of the spreading segment and intruded synrift sediment directly to the north and west.

- 2) Fault patterns and mechanisms responsible for exhuming high and ultrahigh pressure metamorphic rocks from mantle depths ahead of the westward propagating seafloor spreading rift tip, and nature of the rocks above the detachment faults associated with the active D'Entrecasteaux Islands core complexes. Two drilling transects will penetrate sediment and upper plate rocks above the detachment faults, one north of the Prevost Range core complex on Normanby Island (<30 km from the active seafloor spreading tip), and the other north of the Mailolo core complex on northwest Fergusson Island. The active submarine sections of these

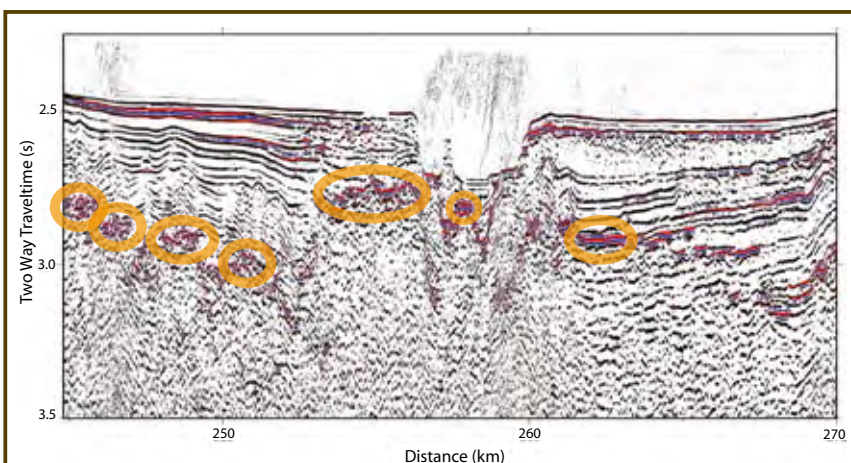


Figure 5. Multichannel seismic profile across the rift graben of the northern Guaymas rift segment of the Gulf of California (see Fig. 4 for location). Sills (circled; identified by reflectivity changes in the sediments they are intruding) drive hydrothermal circulation within the rift valley. Drilling in this vicinity will target thermal and geochemical fluxes as well as geobiological processes including methanogenesis. After Lizarralde et al. (2007).

faults extend below marine sediment that was deposited before and during motion of the hanging wall fault blocks. Drilling will constrain the timing and amount of exhumation directly west of the active seafloor spreading rift and test various models for core complex formation (Abers et al., 2002; Martinez et al., 2001) and for exhumation of high and ultrahigh pressure metamorphic rocks.

North Atlantic Magma-Dominated Margins

Key Aspects and Problems to be Addressed: The conjugate northwest Europe and east Greenland margins (Fig. 7) are characterized by voluminous magmatism associated with the Iceland plume, and their formation may have had significant environmental impact (Eldholm et al., 1989; Eldholm and Grue, 1994; Larsen and Saunders, 1998; Svensen et al., 2004; Storey et al., 2007). The magmatic productivity cannot be explained by simple decompression melting of normal temperature, sub-lithospheric mantle. Three primary competing hypotheses for excessive magmatism are 1) mantle plume with elevated temperatures (White and McKenzie, 1989); 2) small-scale convection at the base of the lithosphere (Mutter et al., 1988; King and Anderson, 1995); and 3) heterogeneities in mantle source composition (Korenaga, 2004).

Regional Setting and Background: The northeast Atlantic conjugate rifted margins show evidence for extensive magmatism including SDRs, igneous intrusions, and high seismic velocity bodies at the base of the crust attributed to magmatic underplating (Fig. 7). The conjugate margins are segmented along strike by the northwest-trending Jan Mayen Fracture Zone and the Bivrost Lineament, which separate the Møre, Vøring, and Lofoten margins and their conjugates at the Jan

Mayen microcontinent and off northeastern Greenland (Eldholm et al., 2002). The margin segments are characterized by different tectono-magmatic styles and sediment distributions, with magmatism decreasing from the Vøring segment to the north and south.

Proposed Drill Sites: The overall plan is for two shallow basement penetration transects (eleven holes) located north and south of the Jan Mayen Fracture Zone (Fig. 7), respectively, and one deep sub-basalt hole within the southern transect (Fig. 7). A dip transect (six holes; Fig. 7) to examine temporal variability of magmatism extends across the conjugate margin segment pair of central Møre/Jan Mayen Ridge. Each hole is well characterized by high quality seismic reflection data. A strike transect to sample

breakup-related volcanic rocks in different margin segments as well as facies units extends along the Norwegian margin. The main segments to be drilled include 1) the central Møre margin (i.e., the location of the dip transect); 2) the southern Vøring margin (transform margin related volcanism); 3) the northern Vøring margin (voluminous volcanic complex); and 4) the southern Lofoten margin (small volcanic complexes). The deep hole—to examine temporal variability of magmatism and the nature, environment, and implications of the rift and early breakup magmatism—will be a reoccupation of a landward site on the Møre-Jan Mayen Ridge conjugate margin transect. Specific issues to be addressed by drilling are 1) melt sources and melting conditions, 2) timing of magmatism, 3) spatial and temporal variations of volcanism, 4) eruption environment and vertical movements, 5) along-axis variations in melt production, and 6) consequences of excessive magmatism for environmental change.

Overall, geophysical (including two-dimensional, three-dimensional, and wide-angle seismic) and geological (including DSDP, ODP, and commercial drilling) data sets for the North Atlantic LIP are comprehensive and of high quality.

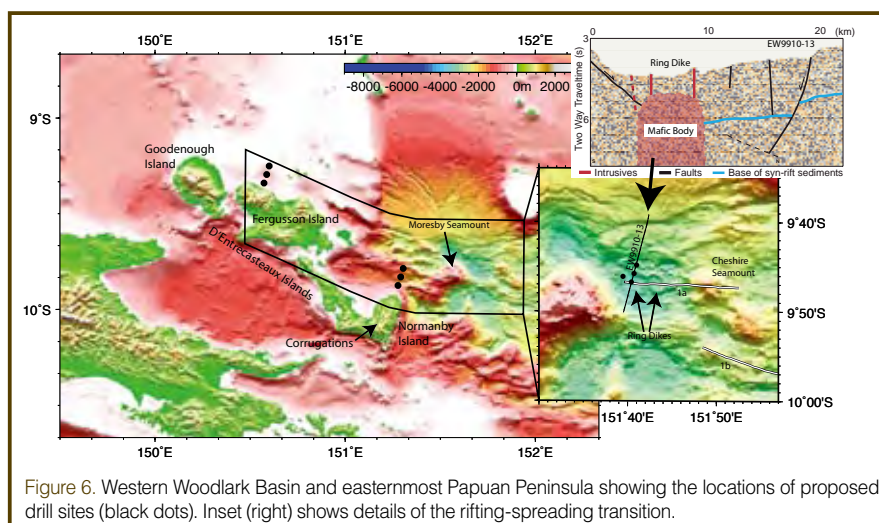


Figure 6. Western Woodlark Basin and easternmost Papuan Peninsula showing the locations of proposed drill sites (black dots). Inset (right) shows details of the rifting-spreading transition.

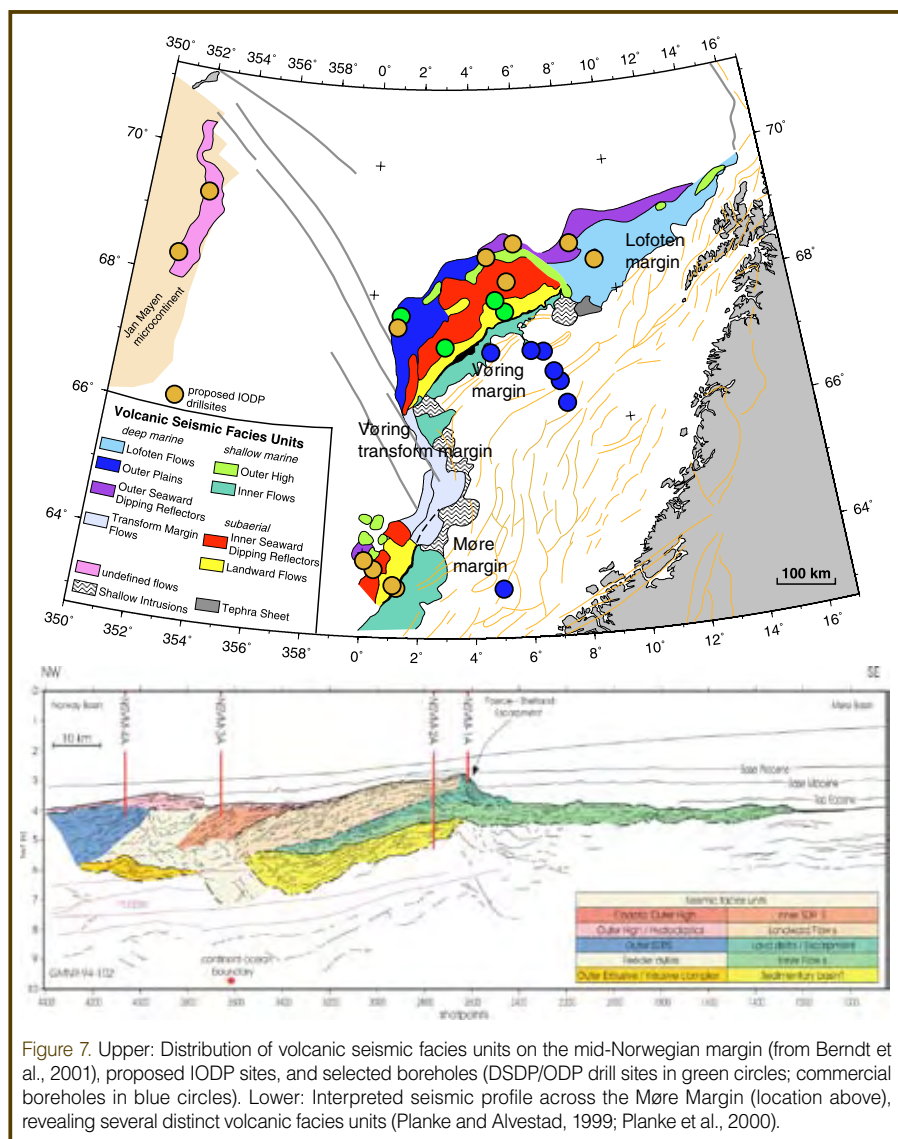


Figure 7. Upper: Distribution of volcanic seismic facies units on the mid-Norwegian margin (from Berndt et al., 2001), proposed IODP sites, and selected boreholes (DSDP/ODP drill sites in green circles; commercial boreholes in blue circles). Lower: Interpreted seismic profile across the Møre Margin (location above), revealing several distinct volcanic facies units (Planke and Alvestad, 1999; Planke et al., 2000).

Newfoundland-Iberia Rift

Key Aspects and Problems to be Addressed: The Iberia and Newfoundland margins (Fig. 8) lack extensive magmatism; they are hyper-extended, characterized by polyphase and diachronous rifting, detachment faulting, and mantle serpentinization and unroofing (Pérez-Gussinyé and Reston, 2001; Reston, 2005; Tucholke et al., 2007). Thin sedimentary cover makes tectonic targets uniquely accessible to drilling. Key problems concern the timing of rifting (along and across the margins), breakup, and the onset of seafloor spreading; the mechanism(s) of extreme crustal thinning; the role of detachment faulting in mantle unroofing; and the nature of basement within the continent-ocean transition.

Regional Setting and Background: During Late Jurassic and Early Cretaceous periods, rifting localized between Newfoundland and Iberia, and the two separated. Breakup propagated from the Central Atlantic northward (Srivastava et al., 2000), but critical details remain controversial. In the south of the Newfoundland-Iberia region, seafloor spreading anomaly M3 (128–130 Ma; Gradstein et al., 2004) has been

identified (Srivastava et al., 2000; Russell and Whitmarsh, 2003), whereas in the north, M0 (124.5–125 Ma) appears to be the oldest anomaly. However, some evidence indicates that pre-spreading rifting continued into late Aptian time (~112 Ma; Boillot et al., 1987; Reston, 2005), supporting the idea that the earliest ‘seafloor spreading’ magnetic anomalies may have originated from unroofed mantle rather than igneous crust formed at a focused spreading center.

Such uncertainties in timing of key events preclude a thorough understanding of dynamic processes because neither rates nor spreading mechanisms are yet accurately known. Another outstanding problem is how the crust was thinned to only a few kilometers, challenging many tenets of lithospheric rheology and isostasy. Some combination of polyphase faulting and DDS seems likely, but the relative importance of the two is controversial (Davis and Kusznir, 2004; Reston, 2007). Similar problems characterize other rifted margins—the South Atlantic (Moulin et al., 2005), Northwest Australian (Driscoll and Karner, 1998), the Labrador Sea (Chian et al., 1995), and the Parentis basin (Pinet et al., 1987).

Proposed Drill Sites: A key objective is to determine the timing of events, which is needed for a quantitative understanding of the rates of processes associated with final thinning, crustal separation, lower crust and mantle exhumation, the onset of mantle melting, and seafloor spreading. Complete sedimentary sections on both margins are required to achieve this objective (Fig. 8). Well defined synrift wedges above a probable detachment (S) on the Iberia (Galicia Bank) margin and relatively thin sedimentary cover on the conjugate Newfoundland (Flemish Cap) margin provide unique opportunities to establish the timing of events. Another major objective is to test competing ideas on how lithosphere deforms during the final thinning phase of extension leading to exhumation of lower crust and mantle. Thinned continental crust must be sampled at key locations to achieve this objective (Fig. 8).

Flemish Cap. Drilling will help determine 1) the role of hypothesized concave-down faults in exhumation of lower crust and upper mantle during late breakup; 2) the interplay among tectonic, magmatic, and serpentinization processes in hyper-extended rift environments; 3) whether continental crust was removed completely amagmatically; 4) if initial melt products were distributed asymmetrically, with more

melt on the Newfoundland side; 5) when and how rifting transitioned to seafloor spreading; 6) controls on the localization of deformation into serpentinized shear zones; and 7) when asymmetries between Galicia Bank and Flemish Cap developed. The most landward sites lie on the continental slope where continental crust is <5 km thick, and will provide key information on rock types where extreme thinning is observed in seismic sections. The sites will sample as much stratigraphy as possible to constrain timing as well as penetrate into basement rocks. The most seaward sites will test competing hypotheses for the formation of transition zone crust, mantle exhumation, and formation of anomalously thin oceanic crust.

Galicia Bank. Drilling will recover complete sedimentary sections at two sites that will help establish the timing and geometry of fault block movements associated with the formation of the 'S' reflection and emplacement of the peridotite ridge just prior to seafloor spreading (Fig. 8). One site (GBB-7A) will penetrate the 'S' reflection, hypothesized to represent a regional detachment surface. Data from overlying strata will constrain the timing of any motion along the surface and dip angles during motion. Coring the hypothesized detachment will reveal deformation above and below the fault while it was active. Another site (GBB-8A) will penetrate a basin and the eastern flank of the peridotite ridge. While the top of the ridge has been sampled, little is known about its internal structure and mode of formation, and when the ridge was exposed at the seafloor.

South Atlantic Margins

Key Aspects and Problems to be Addressed: Large data sets from the conjugate margins of the South Atlantic suggest that the crust here has been thinned more than can be explained by the observed faults. Drilling of well imaged synrift and early postrift sediment infill of marginal sag basins will provide critical information on timing and facies

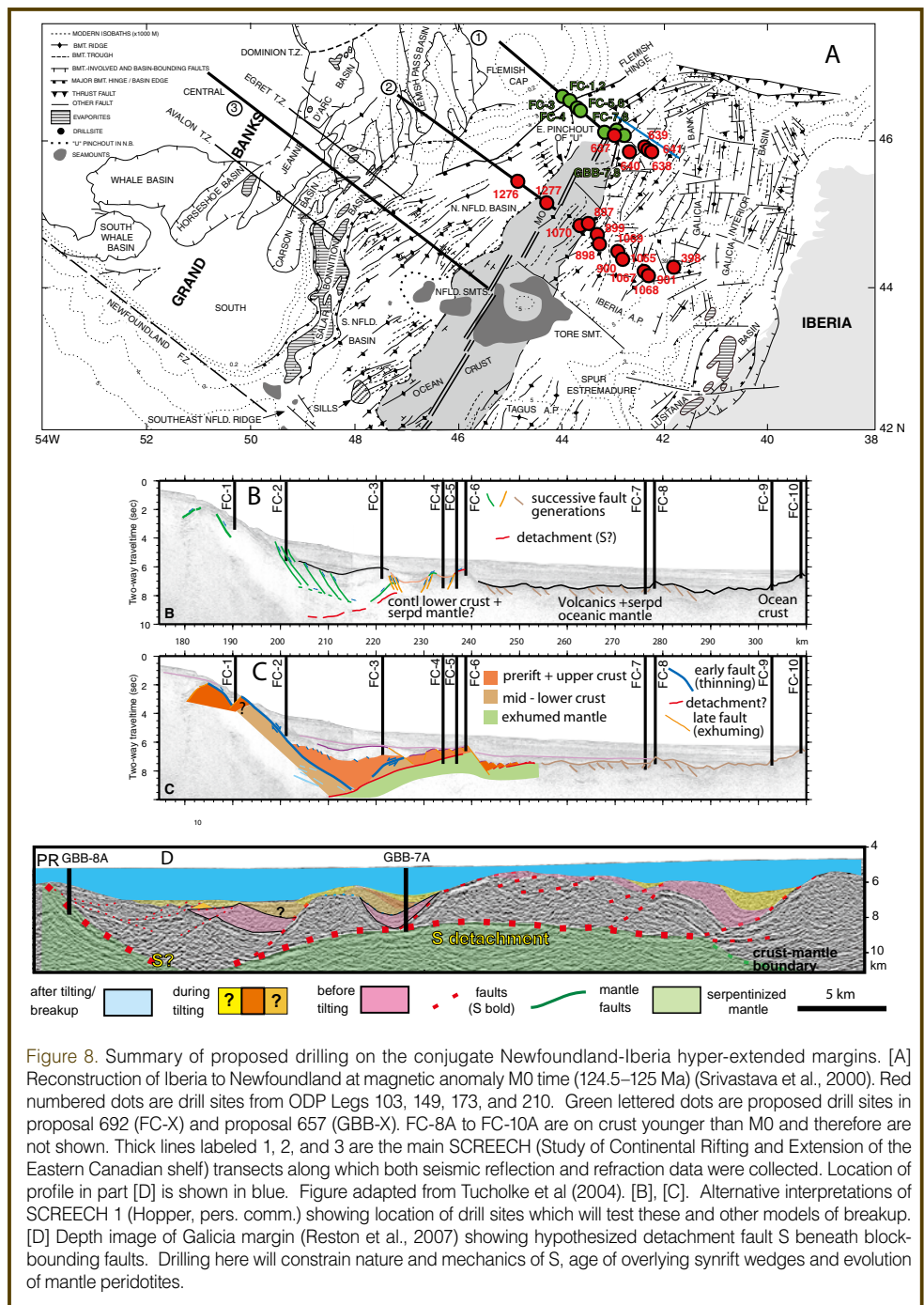


Figure 8. Summary of proposed drilling on the conjugate Newfoundland-Iberia hyper-extended margins. [A] Reconstruction of Iberia to Newfoundland at magnetic anomaly M0 time (124.5–125 Ma) (Srivastava et al., 2000). Red numbered dots are drill sites from ODP Legs 103, 149, 173, and 210. Green lettered dots are proposed drill sites in proposal 692 (FC-X) and proposal 657 (GBB-X). FC-8A to FC-10A are on crust younger than M0 and therefore are not shown. Thick lines labeled 1, 2, and 3 are the main SCREECH (Study of Continental Rifting and Extension of the Eastern Canadian shelf) transects along which both seismic reflection and refraction data were collected. Location of profile in part [D] is shown in blue. Figure adapted from Tucholke et al (2004). [B], [C]. Alternative interpretations of SCREECH 1 (Hopper, pers. comm.) showing location of drill sites which will test these and other models of breakup. [D] Depth image of Galicia margin (Reston et al., 2007) showing hypothesized detachment fault S beneath block-bounding faults. Drilling here will constrain nature and mechanics of S, age of overlying synrift wedges and evolution of mantle peridotites.

for understanding margin evolution, including the cause of the extension discrepancy.

Regional Setting and Background: Rupture between South America and Africa propagated from south to north in the Late Jurassic and Early Cretaceous time as the South Atlantic Ocean formed. The resulting South Atlantic passive margins can be divided broadly into three provinces. The first province, the South of Walvis, voluminous magmatism led to ~100-km-wide SDRs within the crust along both the Argentine and Namibian conjugate margins. The second province, the North of Walvis Ridge—the rifted margin of eastern Brazil and its conjugate Angola, Congo, and Gabon margins (Fig. 9)—also experienced volcanism during breakup, but

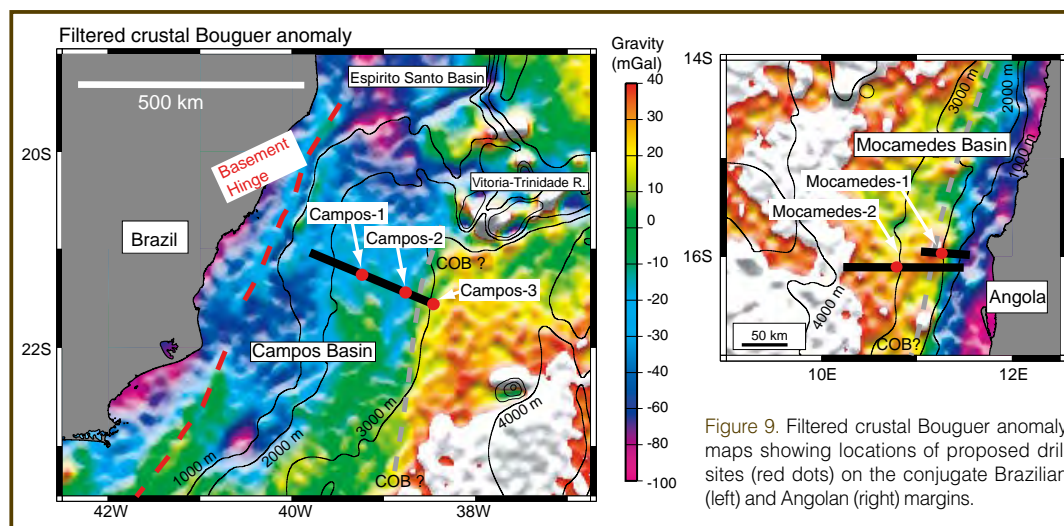


Figure 9. Filtered crustal Bouguer anomaly maps showing locations of proposed drill sites (red dots) on the conjugate Brazilian (left) and Angolan (right) margins.

not sufficiently voluminous to form SDRs. Synrift and postrift sedimentation was dominated by Aptian salt, carbonate platforms, and clastic sediment. The third province, the conjugate rifted margins of the equatorial Atlantic, are narrow compared to the hyper-extended margins to the south. Synextensional sag basins, where salt was deposited and mobilized, are found in both the Campos Basin offshore Brazil and the Kwanza Basin offshore Angola. The relative timing of Early Cretaceous synrift volcanism, evaporite deposition, and onset of seafloor spreading are controversial. The seaward edge of the Aptian salt basin may lie on thick SDR wedges (Jackson et al., 2000), implying that salt was deposited after final continental breakup; volcanic edifices in the continent-ocean transition zone probably acted as barriers between the episodically dry marginal basin and the open ocean where new oceanic crust was forming. However, Aptian salt is also abundant in the vicinity of the Congo Fan, where no thick igneous crust or SDRs are imaged in the continent-ocean transition zone. Even in the absence of a potentially bounding basement high, the seaward limit of autochthonous salt appears to lie close to the inferred landward limit of oceanic crust (Marton et al., 2000). Evaporites may even have been deposited on serpentinized mantle in the continent-ocean transition zone (Moulin et al., 2005). Mantle exhumation prior to continental breakup is well characterized from the Newfoundland-Iberia rift, and conceptual models show how continental mantle may be exhumed by low-angle normal faults (Lavie and Manatschal, 2006). However, whereas mantle was exhumed well below sea level along the Newfoundland rift (Tucholke et al., 2007), the continent-ocean transition zone of the central South Atlantic appears to have been close to sea level prior to breakup.

Proposed Drill Sites: Drill sites in the Campos Basin offshore Brazil and the Moçamedes Basin offshore Angola will establish the tectonic setting of marginal basin formation in the central South Atlantic (Fig. 9). Two sites in the Campos Basin lie at the seaward end of a sag basin; drilling will establish the timing and environment of Aptian salt basin formation. A third site on the Brazilian margin, just seaward of the

characteristics to South Atlantic margins; 2) the age of the first oceanic crust, and relative timing of both continental breakup and the deposition of pre-salt sag sequences; 3) the nature and composition of the crust on which pre-salt sequences were deposited; 4) interpreted exhumed continental mantle in the continent-ocean transition zone; 5) the possible existence of top-basement detachment faults; and 6) synrift and early postrift subsidence along a geophysical transect of the conjugate margins.

NW Australian Magma-Dominated Margin

Key Aspects and Problems to be Addressed: The northwest Australian magma-dominated (rifted) margin (Fig. 10) is segmented, and igneous rock volumes vary considerably along strike, without clear evidence for a related mantle plume (Mutter et al., 1988; Hopper et al., 1992; Symonds et al., 1998; Planke et al., 2000). This makes the margin a strong candidate to test the edge-driven/small-scale convection hypothesis for generating excessive magmatism. Temporal and along-strike variations in melt production and temporal and spatial relationship(s) between rifting and magmatism are secondary objectives.

Regional Setting and Background: The western Australian margin can be divided into four main segments separated by major fracture zones: Argo, Gascoyne, Cuvier, and Perth (Fig. 10). The entire margin exhibits breakup magmatism in the form of sills, SDRs, hyaloclastic buildups, and magmatic underplating that formed during Callovian (~163 Ma) Argo margin breakup and subsequent Valanginian (~138 Ma) Gascoyne, Cuvier, and Perth margin breakup (Planke et al., 2000; Symonds et al., 1998). The conjugate margins have been subducted or obducted. Two main hypotheses—mantle plume (White and McKenzie, 1989) and edge-driven/small-scale convection (Mutter et al., 1988; King and Anderson, 1998; Korenaga, 2004)—have been proposed for the formation of these massive igneous constructions (Coffin et al., 2002; Ingle et al., 2002; Müller et al., 2002). Because it cannot be convincingly tied to a well established hotspot track, the

western Australian margin is a highly promising candidate for testing alternative hypotheses for magmatic margin formation (Planke et al., 2000).

If a plume was involved in the formation of the western Australian margin, then the Wallaby Plateau-Zenith Seamount province probably represents the post-breakup track of the plume. This province extends some 1200 km from the continent, across a continent-ocean transition zone likely associated with composite continental and magmatic plateaus, and into a normal ocean basin, providing an ideal opportunity to examine variations in continental contamination in time and space. Documenting the presence or absence of a geochemical plume signature within igneous basement of the Cuvier margin will therefore be a critical complement to previous and ongoing studies of the North Atlantic igneous province.

Proposed Drill Sites: The objectives are as follows: 1) to distinguish between an edge- vs. plume-driven cause for magmatism along a rifted margin; 2) to examine rift and breakup duration, and subsidence history; 3) to investigate the formation and crustal nature of marginal plateaus; 4) to understand the temporal development of multiple SDR wedges; and 5) to determine age, volume, duration, and environment of volcanism. To address these objectives, a five-hole transect across the southern Cuvier margin and the Wallaby Plateau as well as a reference hole in similar age oceanic crust of the nearby Cuvier Abyssal Plain are envisioned (Figs. 10 and 11). The holes will sample i) multiple SDRs across the margin; ii) the Wallaby Plateau, and iii) oceanic crust. The petrology and geochemistry of the recovered rocks will be used to determine melting conditions, magma reservoir type (asthenosphere, lithosphere, plume), and contamination by continental lithosphere. Precision age determinations will constrain the temporal and spatial evolution of the magma source. Regionally, the west Australian margin is relatively well investigated by seismic surveys, dredging, and commercial drilling on the continental shelf.

Essential Complements to Drilling

To date, seismic studies of rifted continental margins by the academic community have been almost entirely 2-D, comprising widely spaced profiles relative to the lateral scale of faults and stratigraphic variations.

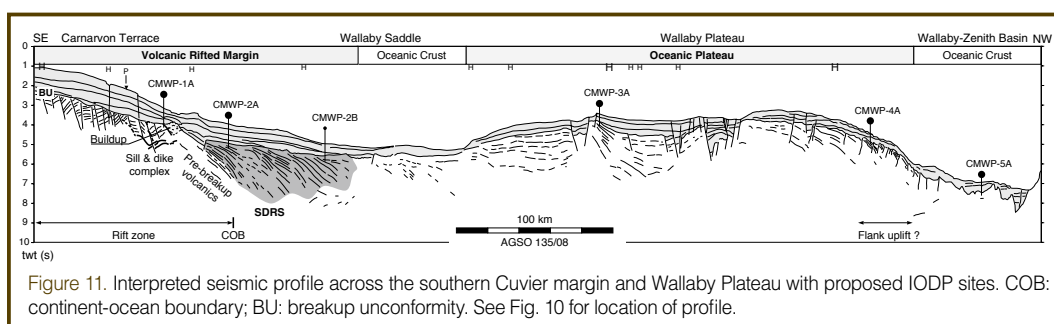


Figure 11. Interpreted seismic profile across the southern Cuvier margin and Wallaby Plateau with proposed IODP sites. COB: continent-ocean boundary; BU: breakup unconformity. See Fig. 10 for location of profile.

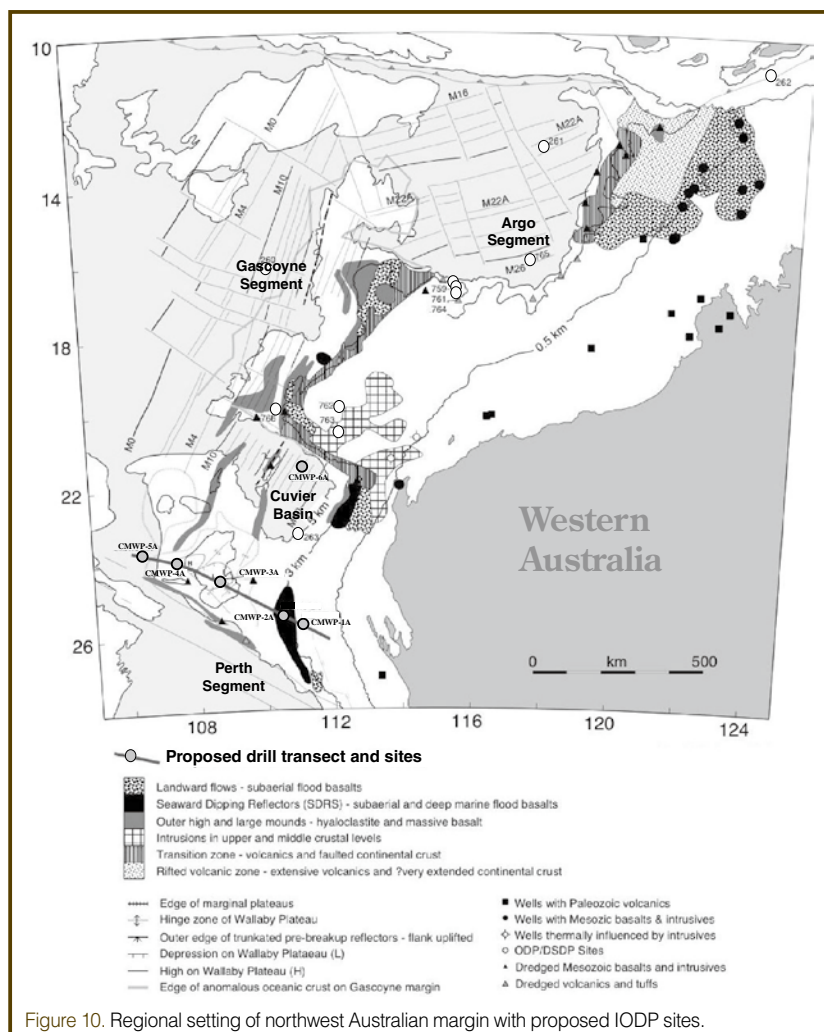


Figure 10. Regional setting of northwest Australian margin with proposed IODP sites.

We now recognize that 2-D seismic technology does not meet the challenge of contemporary research in rifting and breakup, which are fundamentally 3-D processes. As part of its strategy in investigating rifting and breakup, the IODP should integrate acquisition and interpretation of 3-D seismic reflection data with overall program planning.

Land-based geological research has been a valuable complement to ocean drilling and marine geophysics at several rifted margin systems. The Gulf of California is a good example of a rift system that is only partly submerged. Similar opportunities are offered by the Woodlark Basin, where metamorphic core complexes and other rifting components are exposed on islands, and in the North Atlantic LIP, where flood basalts and related rocks crop out in the British Isles



Figure 12. Early Tertiary flood basalt sequence in eastern Greenland associated with rifting and breakup between eastern Greenland and northwestern Europe, highlighting opportunities along some margins for synergetic onshore-offshore studies.

and Greenland (Fig. 12). The study of rift systems now exposed on land in mountain belts is another way that the advantages of field geology can contribute to the study of rift systems, as exemplified by comparisons between the Alpine Tethys rift system exposed in the Alps (Fig. 13) and the Iberia/Newfoundland rift system (Manatschal, 2004; Lavier and Manatschal, 2006).

Modeling the processes of continental rifting is increasingly important for making predictions that can be tested by drilling. Forward dynamic models now provide insights into magmatism accompanying breakup (Boutlier and Keen, 1999; Nielsen and Hopper, 2004) as well as primary controls on passive margin width, symmetries/asymmetries, and evolution from wide, diffuse rifting into narrow, localized rifting resulting in the formation of passive margins (Huisman and Beaumont, 2003, 2007; Lavier and Manatschal, 2006).

The study of continental breakup and sedimentary basin formation offers opportunities for significant collaboration between the hydrocarbon industry and the IODP. During the

workshop, interest in collaboration was highest for the conjugate rifted margins of the South Atlantic.

Technical Requirements

COBBOOM will require a combination of existing IODP technology and the development of new technology. Proposed deep and challenging drilling will require the use of both riser and non-riser drills. Where water depths are less than 2500 m, we anticipate using riser capability, casing, and mud circulation to increase hole stability and improve the likelihood of deep penetration. However, some holes have the goal of sampling highly stretched continental crust and upper mantle beneath relatively thin sediment cover and in much deeper water. For such holes, ultra-long drill strings deployed in a non-riser mode and supported by extensive casing programs will be required.

Improved methods of core orientation in sedimentary and crystalline rocks will improve our ability to relate microstructures and faults in the core to the strain distribution in the rifting system. In active rift environments, borehole observatories will be used to monitor the presence and pressure of fluids in faults active during rifting. Technology currently being developed to monitor microseismicity near the boreholes will assist in understanding its relationship to fluid pressure variations.

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exhumed upper mantle in the Alps. Finally, we acknowledge the essential contributions of Kelly Kryc of IODP-MI who organized the workshop.

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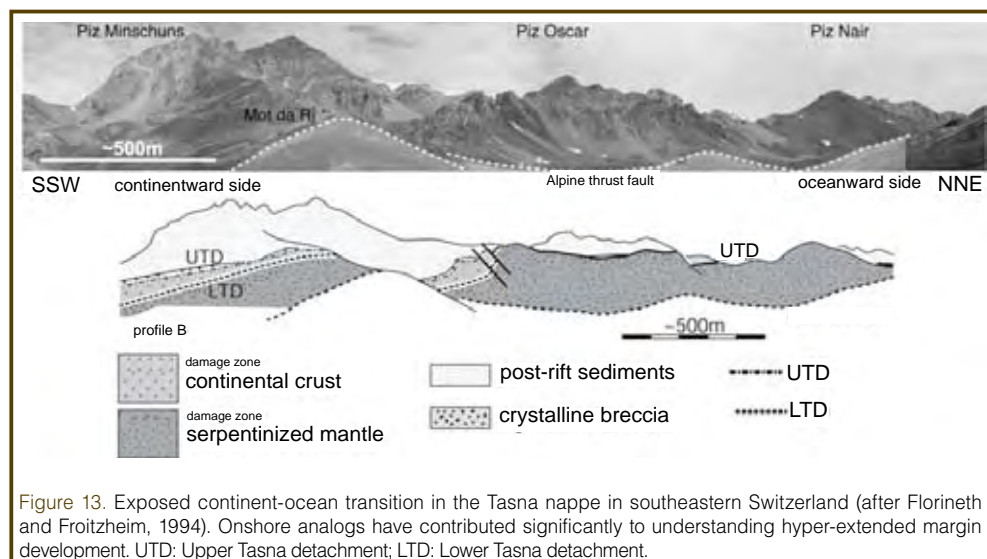


Figure 13. Exposed continent-ocean transition in the Tasna nappe in southeastern Switzerland (after Florineth and Froitzheim, 1994). Onshore analogs have contributed significantly to understanding hyper-extended margin development. UTD: Upper Tasna detachment; LTD: Lower Tasna detachment.

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