

Monitoring of Rock Mass Behavior at the Closest Proximity to Hypocenters in South African Gold Mines

by Hiroshi Ogasawara and the Research Group for Semi-controlled Earthquake-Generation Experiments in South African Deep Gold Mines

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To bridge the scale gap between a laboratory scale and a large earthquake scale, we need experiences in an intermediate scale at the smallest distance from the hypocenters (Fig. 1). However, the closest monitoring is not always easy, and the San Andreas Fault Observatory at Depth (SAFOD) is an exceptional opportunity to obtain experiences from natural small earthquakes (Hickman et al., 2004). Tremors in deep mines resemble natural small earthquakes, being also good to take a close look at the hypocenters (Spottiswoode and McGarr, 1975). Recently, modern monitoring is carried out with wider dynamic range and frequency band (Mendecki, 1997). The Research Group for the Semi-controlled Earthquake-Generation Experiments in South African Deep Gold Mines (SeeSA) have been co-operating with ISS International Ltd. and South African gold mines to study rock mass behavior (Iio, 1995; Iio and Fukao, 1992; Ishii and the Research Group for Semi-Controlled Earthquake Generation Experiment at South African Gold Mines, 1996; Nicolaysen, 1992; Ogasawara et al., 2001, 2002a, 2005a, 2005b; Sumitomo, 1998). We have deployed eight experimental sites at four gold mines at depths of a few km (Fig. 2; Ogasawara et al., 2002a, 2005a, 2005b), and at one of these we are collaborating with ICDP-DAFSAM/NSF-NELSAM project (PI: Ze'ev Reches).

In 1996, at the Western Deep Levels South Mine (currently Mponeng mine; Fig. 2), we deployed an array of nine triaxial accelerometers within 100 m of potential hypocenters of M>2 earthquakes, successfully delineating the finest details of small earthquakes. The relationship between seismic moment and corner frequency was the same as that for

medium or large natural earthquakes (Ogasawara et al., 2002b). The source processes of some M-1 or smaller earthquakes were as complicated as large, natural earthquakes (Yamada et al., 2005).

During 2000–2003 at the Bambanani mine (Figs. 2 and 3), we successfully recorded the rock-mass behavior in the entire life span of M2.4 and M2.5 earthquakes (Mw2.9 and Mw2.0, respectively, in Fig. 4a; Ishii et al., 2000; Ogasawara et al., 2005a; Takeuchi, 2005). We installed an Ishii strainmeter (Ishii et al., 1997) within 100 m of the seismic sources and recorded the data continuously with resolutions of 25-Hz and 24-bit. The initial significant change was by an Mw0.2 event in September 2001 within a few tens of meters. The largest event (Mw3.0) within 250 m of the strainmeter took place in 2002, however, not close enough to cause the largest strain change. The largest strain changes were associated with two M2 events within 100 m of the strainmeter in 2003 (yellow circles in Fig.3, upper). The strain change was as large as 100 microstrain, corresponding to several MPa. Such a large change is recorded only at the closest distance from the seismic fault. Channel 1 subparallel to the maximum principal stress was contracting until the M2 sequence and then was released by the sequence.

However no accelerating strain preceded the hundreds of co-seismic steps for the catalogued seismic events, even for those of ~100 microstrain (stress drop ~7 MPa) (Ogasawara et al., 2005a; Takeuchi, 2005).

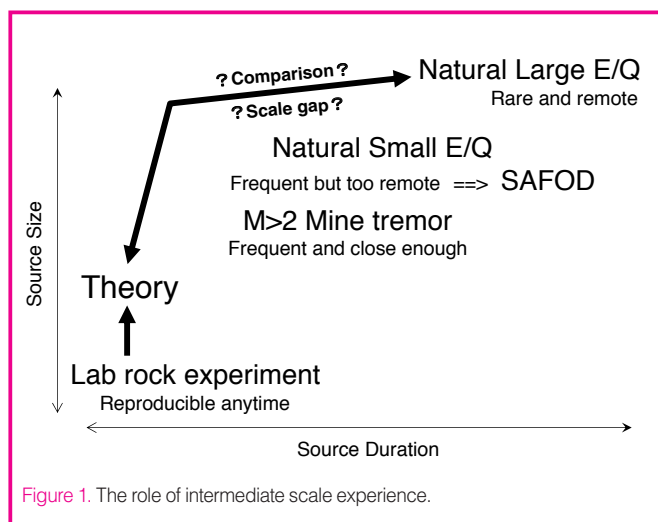


Figure 1. The role of intermediate scale experience.

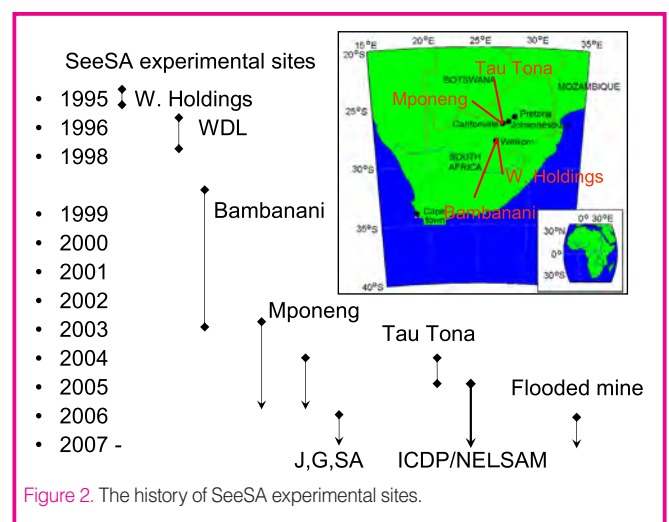


Figure 2. The history of SeeSA experimental sites.

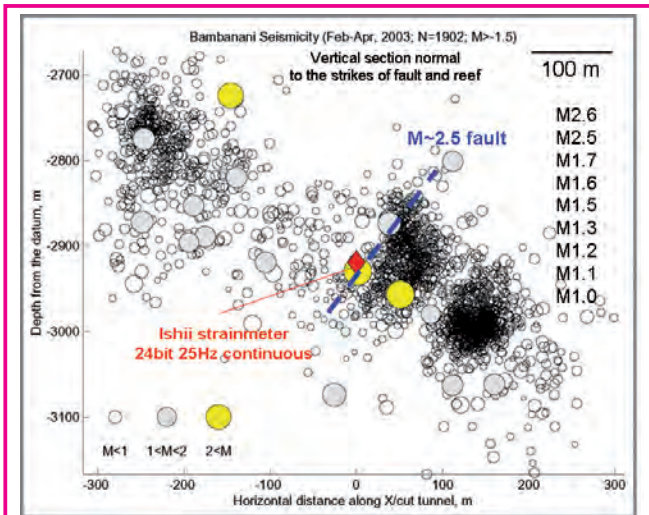


Figure 3. Configuration of an Ishii strainmeter, the M2.5 fault and the seismicity (Feb-Apr 2003) at the site at the Bamabanani mine, Welkom.

Thoroughly going through the strain recordings, Yamamoto et al. (2006) found that the Ishii strainmeter picks up seismic events smaller than $M = -1$ (much more than those catalogued by the mine's seismic network). They also found hundreds of the aftershocks having variable senses of strain steps and postseismic changes. This suggests the strainmeter was located within the aftershock area.

The step response to the catalogued earthquake accompanied overshoot and ringing (Fig. 4b; hereinafter, **dynamic**

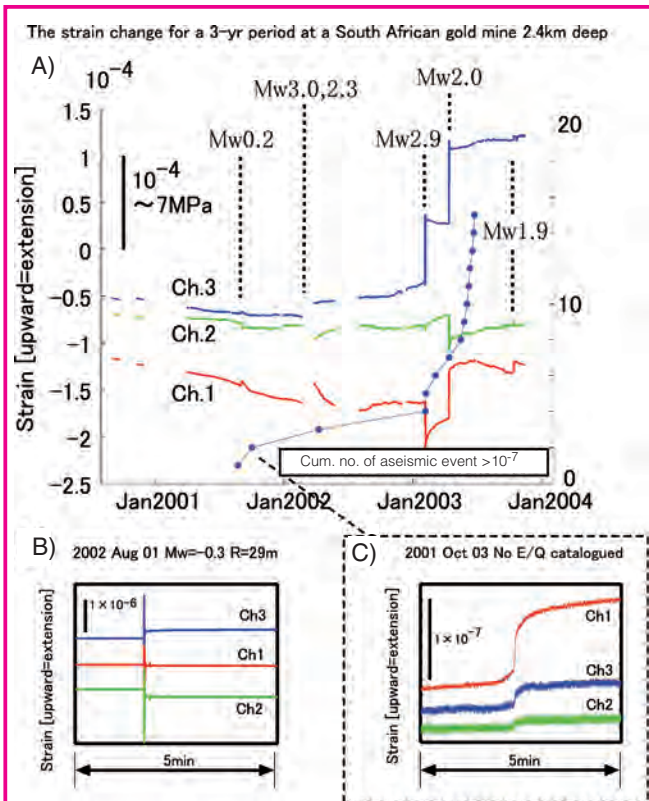


Figure 4. [A] A three-year strain changes (red, green, and blue lines) and cumulative number of aseismic events (LDR events; connected blue dots). [B] A typical response to a catalogued seismic event. Note the dynamic response (overshoot and ringing) associated with the strain step. [C] A typical example of slow step with clear forerunner (after Naoi et al., 2006 b).

Statistics of strain events $> 10^{-7}$

Out of 70 events ($> 10^{-7}$)

{ 38: catalogued E/Q ($M > -0.5$)

{ 32: others

{ 6: Significant Dynamic Response

{ 11: Not easily categorized

{ 15: Little Dynamic Response (LDR)

{ 3: Very slow, some preceded by forerunner

{ 12: Others

Figure 5. The statistics of the strain events (data after Naoi et al., 2006a).

response). However, Naoi et al. (2006a) found that slow strain-steps as small as 0.1 microstrain were also frequently recorded with little dynamic response (LDR). They found that the step durations were variable and significantly longer than normal seismic steps (Fig. 4c; the moments possibly correspond to $M = -1$ or smaller seismic events). Interestingly, most LDR events took place after the large strain change and few very slow events were preceded by the significant accelerating strain (Figs. 4c and 5).

The entire lifespan of a large natural earthquake includes loading, foreshocks, main rupture, aftershocks, and postseismic deformation. Slow events are also associated. The forerunning, accelerating deformation is not often observed, however.

For $-1 < M < 2$ mine tremors, we again observed almost the same set of phenomena. Additionally, we could find the clear forerunners associated only with the small, slow strain event. We were able to see them because we were able to install strainmeters within seismic sources.

Monitoring the smaller subsets of the phenomena (gray area in Fig. 6) could be interesting. To get the time and magnitude range closer to those for AE, we start the strain

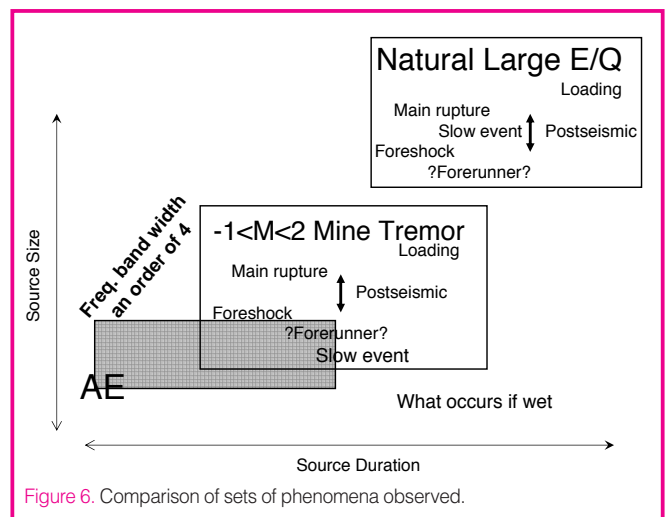


Figure 6. Comparison of sets of phenomena observed.

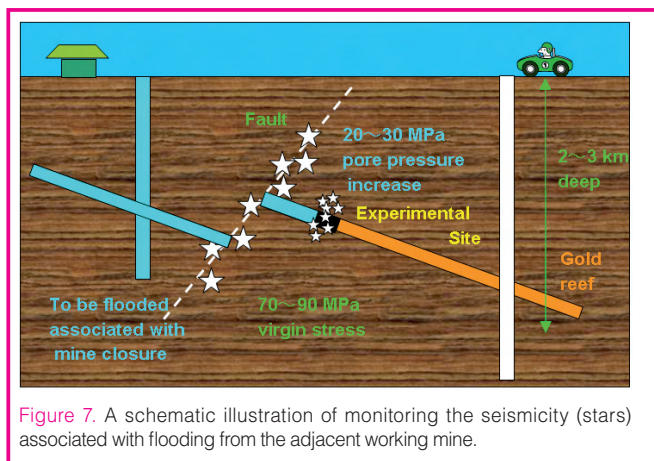


Figure 7. A schematic illustration of monitoring the seismicity (stars) associated with flooding from the adjacent working mine.

monitoring with much higher sampling rate at new sites. One of these is carried out collaterally with Nakatani et al., (2006) who attempted to monitor the wide dynamic/frequency range of fracturing process from AE (up to 200 kHz) to seismic events at a potential seismic fault.

At the previous experimental sites, the rock mass was not saturated with water. We also have to learn what behavior can be seen under wet conditions, and along these lines, the flooded South African gold mines are very unique cases to provide us with good experimental sites. We can investigate 1) the effects of flooding and the corresponding rising water levels on the stability of faults and other geological features, 2) the effects of seismicity on inter-mine water plugs and mine barriers pillars, and 3) seismic damage risks to neighboring mines in areas in which mines are mature.

These investigations are very crucial because the behavior of highly stressed rock mass during flooding is presumably analogous to that during natural earthquake swarms or stable or unstable slip at natural great earthquake hypocenters under water-saturated conditions. We cannot access inside the earthquake swarm in Japan, but we can do it in South Africa.

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